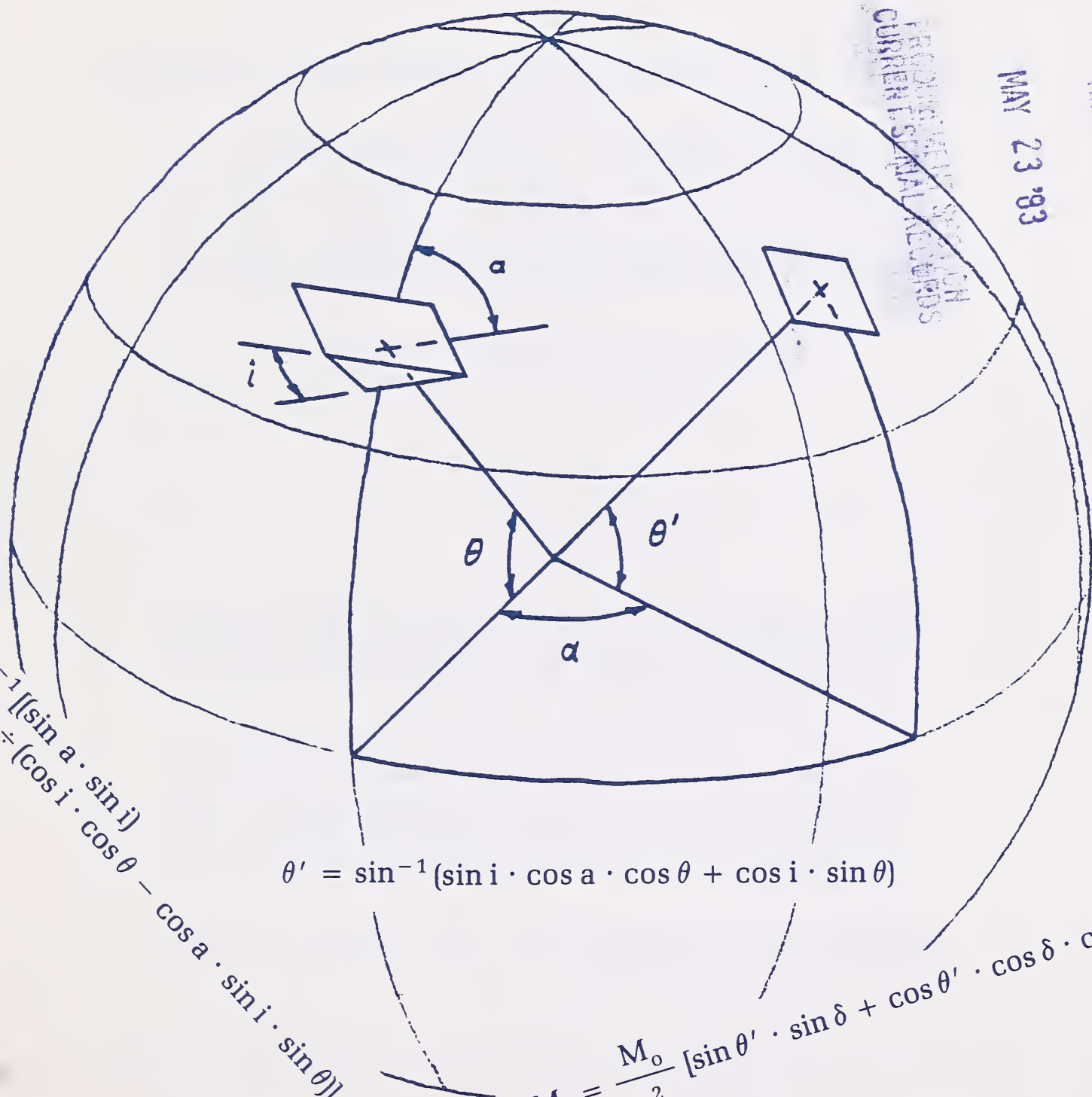


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Determination of Potential Direct Beam Solar Irradiance

Merrill R. Kaufmann and James D. Weathered



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Abstract

Procedures are presented for calculating potential direct beam solar irradiance, corrected for latitude, azimuth and inclination of slope, date, and time of day. Equations are structured to permit the user to calculate instantaneous or total daily irradiance using total incoming shortwave irradiance or any selected portion of the solar spectrum.

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Keywords: Solar irradiance, direct beam irradiance, latitude, declination, azimuth, inclination

Determination of Potential Direct Beam Solar Irradiance

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Determination of Potential Direct Beam Solar Irradiance

Merrill R. Kaufmann and James D. Weatherred

Introduction

Potential direct beam solar irradiance is the radiation received above the earth's atmosphere. The maximum direct beam irradiance is received when a surface above the atmosphere is perpendicular to the sun's rays. For incoming shortwave irradiance, this is equivalent to the solar constant, generally taken to be $1360 \text{ W} \cdot \text{m}^{-2}$ ($1.95 \text{ cal} \cdot \text{cm}^{-2} \cdot \text{min}^{-1}$).

Actual irradiance received at the earth's surface is lower than the potential direct beam irradiance. The reduction results from a number of factors, some of which are highly predictable because they are geometric in nature, and some of which are less predictable because of varying atmospheric effects.

Geometric effects include latitude, solar declination, distance between the sun and earth (radius vector), inclination and aspect of the plane in question, time of day, and obstructions. The plane or surface may be a hillside, leaf, wall or roof of a building, solar collector, etc. Atmospheric effects include depth of clear atmosphere through which radiation must pass (a function of solar declination, latitude, elevation, and time of day), radiation scattering (a function of atmospheric composition and contamination), and cloudiness (Gates 1980).

While knowledge of the actual irradiance received at a surface is useful to hydrologists, plant physiologists, solar energy engineers, and others, the complexities of predicting actual irradiance are great because atmospheric conditions are dynamic. Knowledge of potential direct beam irradiance is useful because potential irradiance is a maximum limit for actual irradiance and is usually used to derive estimates of actual irradiance.

Incoming shortwave irradiance is of interest to hydrologists, micrometeorologists, and solar engineers because of its direct role in determining the energy balance of a given site. In some cases, however, only a portion of the solar spectrum is of concern. For example, a plant physiologist studying photosynthesis or stomatal behavior may be concerned only with photosynthetic photon flux density (visible irradiance, 400 to 700 nm), while a human pathologist evaluating skin disorders may be interested in the ultraviolet portion of the spectrum.

Frank and Lee (1966) provided tables giving the times of sunrise and sunset and the total daily incoming irradiance, which covered a range of latitudes, aspects, slopes, and times of year. Buffo et al. (1972) also provided extensive tables and figures of direct beam solar irradiance. Swift (1976) presented an algorithm for calculating daily total solar irradiance on mountain slopes. This algorithm is suitable for use as a computer subroutine, making the use of tables unneces-

sary. Formulas for calculating direct beam irradiance presented by Frank and Lee (1966), Gates (1980), and Lee (1978) are useful, but they require more extensive development by the user for routine prediction of irradiance. Furthermore, the formulas are considered mainly in the context of total incoming shortwave irradiance.

This paper has two objectives. First, the formulas used by Frank and Lee (1966), Buffo et al. (1972), and Swift (1976) are extended to provide more details of the calculation procedures, particularly with regard to determinations of hour angles for calculating sunrise, sunset, and radiation and of the effects of intervening obstruction by terrain. This is done to facilitate the use of computers for calculations of both instantaneous and total direct beam irradiance. The reader is referred to the programs listed in the appendixes and to Swift's (1976) algorithm for linking the calculations.³

Second, the formulas are presented as a general case for direct beam solar irradiance, where the incoming irradiance may be total shortwave, visible, ultraviolet, etc., as selected by the user. In effect, the formulas for instantaneous and total irradiance yield a "multiplication factor" to be taken times the potential irradiance (e.g., the solar constant for total solar shortwave irradiance, approximately $2,600 \mu\text{E} \cdot \text{m}^{-2} \cdot \text{sec}^{-1}$ for 400 to 700 nm irradiance, etc.). Duffie and Beckman (1974) provide data from several sources describing the standard spectral distribution of extraterrestrial irradiance. It must be recognized that the calculation procedures given below pertain only to the determination of irradiance above the atmosphere. Atmospheric effects on actual irradiance received at a surface in question must be determined for the appropriate portion of the solar spectrum of interest.

Abbreviations

M_o	potential irradiance constant (given a dimensionless value of 1.0 for a surface above the atmosphere, perpendicular to the sun's rays, at the equinox)
R_o	potential irradiance constant in user's units
M_i	multiplier for instantaneous irradiance for a given latitude, aspect, inclination, time of year, and time of day
R_i	instantaneous irradiance in user's units
M_t	multiplier for total daily irradiance from sunrise to sunset for a given latitude, aspect, inclination, and time of year

³For those who have access to a Hewlett Packard 41C hand calculator, a Users' Library Solutions Manual for Solar Engineering (No. 00041-90138) has a program for solar-beam irradiation.

R_t	total daily irradiance in user's units
r	radius vector, the ratio of the distance between the sun and earth at a given date to the mean distance
n	Julian date
m	month of year (1 to 12)
d	day of month
θ	latitude in degrees of the given slope, positive in northern hemisphere, negative in southern hemisphere
θ'	latitude of the equivalent slope
δ	solar declination in degrees for the given slope
α	change in hour angle in degrees from the given to the equivalent slope
ω	angular velocity of the earth's rotation ($15^\circ \cdot \text{hr}^{-1}$)
t	time in hours from solar noon
t_1	sunrise at the given slope
t_2	sunset at the given slope
t'_1	sunrise at the equivalent slope
t'_2	sunset at the equivalent slope
t''	time at the equivalent slope for calculating instantaneous irradiance
a	azimuth of the slope in degrees measured clockwise from north
i	inclination of the slope, 0° to 90°
A	solar altitude in degrees
A'	solar azimuth in degrees measured clockwise from south

Equations

Horizontal Surfaces

The basic equation for calculating instantaneous irradiance on a horizontal surface outside the atmosphere is given by Frank and Lee (1966):

$$R_i = \frac{R_o}{r^2} (\sin \theta \cdot \sin \delta + \cos \theta \cdot \cos \delta \cdot \cos \omega t). \quad [1]$$

The value of r^2 ranges from 0.96676 on January 3 to 1.03370 on July 5 (List 1971, table 169). As an approximation, r^2 can be estimated by first determining the Julian date and declination as follows:

$$n = 31(m - 1) + d - 0.4m - 1.8. \quad [2]$$

Round n to the nearest day. If m is 1 (January), add 2 to n . If m is 2 (February), add 3 to n . In leap years, add 1 day if m is 3 or greater (Ball 1978).

Solar declination can be estimated:

$$\delta = 23.5 \cdot \sin [0.9863(284 + n)]. \quad [3]$$

Finally, r^2 is calculated:

$$r^2 = 0.999847 + 0.001406(\delta). \quad [4]$$

By these procedures, δ is estimated within $+0.73^\circ$ to -1.38° , and r^2 is estimated within $+0.84\%$ to -0.70% , depending upon time of year.

Multipliers for instantaneous irradiance may be calculated with a formula similar to equation 1:

$$M_i = \frac{M_o}{r^2} (\sin \theta \cdot \sin \delta + \cos \theta \cdot \cos \delta \cdot \cos \omega t) \quad [5]$$

where M_o is assigned a value of 1.0 at the equinox. In later illustrations, values of M_i are plotted as a function of several variables. The user may calculate instantaneous irradiance using a selected R_o as follows:

$$R_i = M_i R_o. \quad [6]$$

For a horizontal surface, if the sum of the absolute values of θ and δ is less than or equal to 90° , the hour angles of sunrise ($-\omega t$) and sunset (ωt) before and after solar noon can be found as follows:

$$\omega t = \cos^{-1} (-\tan \theta \cdot \tan \delta). \quad [7]$$

However, if the sum of the absolute values of θ and δ is greater than 90° , then the sun remains either above or below the horizon all day. If θ and δ have the same sign, the sun remains above the horizon and ωt equals 180° . If θ and δ have different signs, the sun is below the horizon for the day and $\omega t = 0^\circ$.

Times for sunrise (t_1) and sunset (t_2) in hours before or after mean true solar noon are given by:

$$t_2 = -t_1 = \omega t / 15. \quad [8]$$

The solar times calculated above are mean solar times. True solar time differs from mean solar time by an amount that varies through the year. True solar time may be calculated by algebraically adding a correction, called the equation of time, which ranges from -14 to $+16$ minutes (List 1971, table 169). Frank and Lee (1966) give procedures for converting solar time to local standard time.

Multipliers for total irradiance for a solar day may be calculated by integrating equation [5]:

$$M_t = (M_o / r^2) [2t \cdot \sin \theta \cdot \sin \delta + (12/\pi) \cdot \cos \theta \cdot \cos \delta \cdot 2 \cdot \sin \omega t]. \quad [9]$$

The user may calculate total irradiance using a selected R_o as follows:

$$R_t = M_t \cdot R_o. \quad [10]$$

Note that M_t has time units of hours; therefore R_o should have units of hours.

An example set of calculations for a horizontal surface is given in the Appendix (example 1).

Tilted Surfaces

When a surface is not horizontal, an "equivalent" horizontal surface exists at a higher or lower latitude and longitude which is parallel to the given surface. This equivalent surface is used to calculate the instantaneous and total irradiance received on the given surface. The next series of formulas, extending those of Frank and Lee (1966, citing earlier references), determine the location of the equivalent surface. Refer to figure 1 for a depiction of a given and equivalent slope.

The latitude of the equivalent surface is determined as follows:

$$\theta' = \sin^{-1}(\sin i \cdot \cos a \cdot \cos \theta + \cos i \cdot \sin \theta) \quad [11]$$

where inclination is in degrees and azimuth is 0° to 360° . If inclination is given in percent slope, then:

$$i = \tan^{-1}(\% \text{ slope}/100). \quad [12]$$

The longitudinal correction of hour angle for the equivalent surface is given by:

$$\alpha = \tan^{-1} \left[\frac{(\sin a \cdot \sin i)}{(\cos i \cdot \cos \theta - \cos a \cdot \sin i \cdot \sin \theta)} \right] \quad [13]$$

when a is greater than 0° .

Because the \tan^{-1} of an angle is defined for angles between -90° and 90° , if the slope is east-facing and α is less than 0° , then α is added to 180° . If the slope is west-facing and α is greater than 0° , α is added to -180° . If a is 0° (due north) and the sum of $|\theta|$ and i is

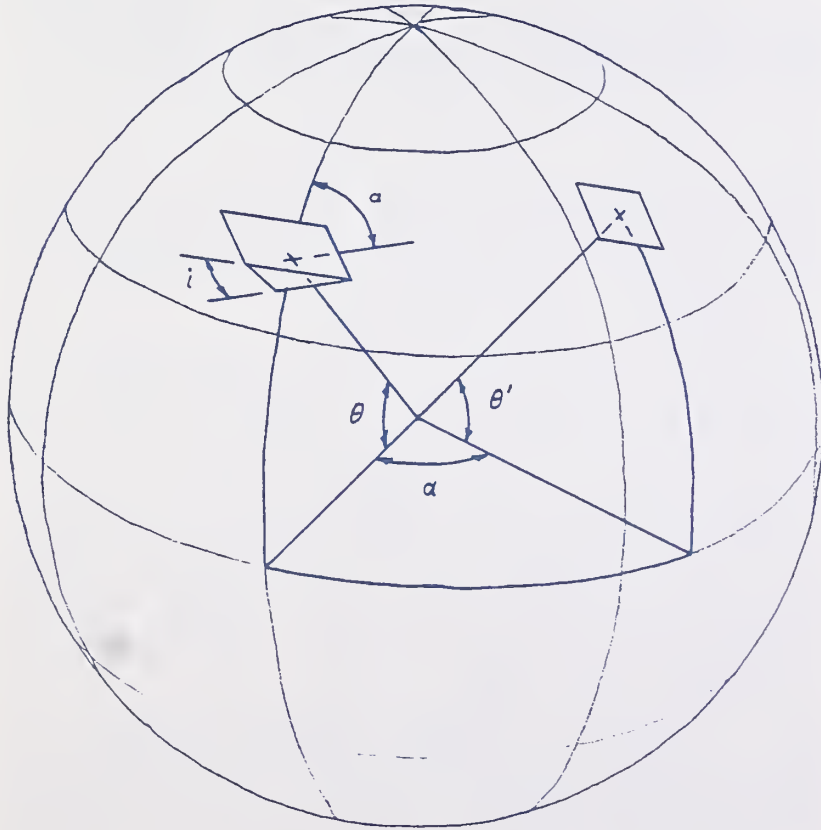


Figure 1.—Relation of given slope and equivalent slope. Given slope at latitude θ has azimuth a and inclination i . Equivalent slope has a different latitude, θ' , and the longitude differs from that of the given slope by the hour angle, α .

less than or equal to 90° , α equals 0° . However, if the sum of $|\theta|$ and i is greater than 90° , α equals 180° .

Hour angles for sunrise and sunset at the equivalent surface are calculated the same way as those for horizontal slopes in equation [7], except that the equivalent latitude is used:

$$\omega t' = \cos^{-1}(-\tan \theta' \cdot \tan \delta). \quad [14]$$

The same limits apply to $(-\tan \theta' \cdot \tan \delta)$ and $\omega t'$ using θ' as to $(-\tan \theta \cdot \tan \delta)$ and ωt for equation [7]. The special case of a double sunrise and sunset is considered later.

Hour angles for sunrise and sunset of the given slope may be determined as follows. If $(-\omega t + \alpha)$ is greater than or equal to $\omega t'$ (values from eq. [7], [13], and [14]), then the given slope is completely shaded throughout the day (t_1 and t_2 equal 0). If the slope is not completely shaded, the time of sunrise at the given slope then becomes:

$$t_1 = [\text{maximum of } (-\omega t) \text{ and } (-\omega t' - \alpha)] / 15. \quad [15]$$

The time of sunset at the given slope becomes:

$$t_2 = [\text{minimum of } (\omega t) \text{ and } (\omega t' - \alpha)] / 15. \quad [16]$$

Irradiance calculations for the equivalent surface require sunrise and sunset times for the equivalent surface. Sunrise for the equivalent slope is:

$$t_1' = [\text{maximum of } (-\omega t + \alpha) \text{ and } (-\omega t')] / 15. \quad [17]$$

Sunset for the equivalent slope is:

$$t_2' = [\text{minimum of } (\omega t + \alpha) \text{ and } (\omega t')] / 15. \quad [18]$$

Finally, irradiance multipliers can be calculated. Instantaneous irradiance is determined as follows:

$$M_i = \frac{M_o}{r^2} [\sin \theta' \cdot \sin \delta + \cos \theta' \cdot \cos \delta \cdot \cos (\omega t'')] \quad [19]$$

where $\omega t'' = \omega t + \alpha$ and t is greater than or equal to t_1 and less than or equal to t_2 . If t is less than t_1 or greater than t_2 , then M_i equals 0.

A series of graphs discussed below present M_i as a function of latitude, azimuth, inclination, date, and time of day. Recalling equation [6], instantaneous irradiance in user units (R_i) can be determined as the product of M_i and R_o . Alternatively, the numerator value of M_o in equation [19] may be replaced with R_o .

Multipliers for total daily irradiance are calculated:

$$M_t = (M_o/r^2) [(t_2' - t_1') \cdot \sin \theta' \cdot \sin \delta + (12/\pi) \cdot \cos \theta' \cdot \cos \delta \cdot (\sin \omega t_2' - \sin \omega t_1')]. \quad [20]$$

Total irradiance in user units (R_t) is determined as the product of M_t and R_o , or R_t may be determined by

substituting R_0 for the numerator value of M_0 in equation [20]. Values of R_t , using the solar constant as R_0 , for a range of latitudes, azimuths, inclinations, and dates are tabulated in Frank and Lee (1966).

An example set of computations for a tilted surface is given in the Appendix (example 2).

Double Sunrise and Sunset

Steep, north-facing slopes and the north sides of buildings at middle latitudes in the northern hemisphere exhibit a sunrise and sunset both in the morning and in the afternoon. Double sunrises and sunsets occur when the following conditions are met:

$$(-\omega t') < (\omega t + \alpha - 360) < (-\omega t + \alpha) < (\omega t'). \quad [21]$$

If these conditions exist, sunrise and sunset times for the given slope are calculated as follows:

$$\text{First } t_1 = -\omega t / 15 \quad [22]$$

$$\text{First } t_2 = (\omega t' - \alpha) / 15 \quad [23]$$

$$\text{Second } t_1 = (360 - \omega t' - \alpha) / 15 \quad [24]$$

$$\text{Second } t_2 = \omega t / 15. \quad [25]$$

For calculating irradiance at the equivalent surface, sunrise and sunset times are:

$$\text{First } t'_1 = -\omega t' / 15 \quad [26]$$

$$\text{First } t'_2 = (\omega t + \alpha - 360) / 15 \quad [27]$$

$$\text{Second } t'_1 = (-\omega t + \alpha) / 15 \quad [28]$$

$$\text{Second } t'_2 = \omega t' / 15. \quad [29]$$

Calculation of the total daily irradiance multiplier for days having two sunrises and sunsets requires determining M_t using equation [20] for the first solar day (first t'_1 to t'_2) and for the second solar day (second t'_1 to t'_2). These values are added to obtain the total daily irradiance multiplier. Again, total daily irradiance in user's units may be determined as the product of M_t and R_0 or by replacing the numerator value of M_0 with R_0 in equation [20].

An example set of calculations for double sunrise and double sunset times is given in the Appendix (example 3). For radiation calculations on days of double sunrise and sunset, refer to example 2.

Obstruction of the Horizon

In mountainous regions, few sites are characterized by a direct view of the earth's horizon. Unless sites are located upon the highest ridgetops, intervening hills and ridges obscure the sun, resulting in later sunrises or earlier sunsets than those predicted by equation [8]. The delay of sunrise, advancement of sunset, or midday obstruction by intervening terrain (or objects) can be determined by comparing the elevation angle of the obstruction with the solar altitude. When the solar altitude at a given azimuth is below the obstruction altitude at the same azimuth, irradiance is set to zero.

Procedures for determining the solar altitude and azimuth are given by Lee (1978). The solar altitude is calculated as follows:

$$A = \sin^{-1}(\sin \theta \cdot \sin \delta + \cos \theta \cdot \cos \delta \cdot \cos \omega t). \quad [30]$$

Values of A less than 0° indicate that the sun is below the observer's true horizon, while values greater than 0° indicate that the sun is above the observer's true horizon.

The computation of solar azimuth is somewhat more difficult. Solar azimuths calculated below are determined clockwise from the observer's south. If $|\theta|$ is greater than $|\delta|$, the solar azimuth is given by

$$A' = \sin^{-1}(\cos \delta \cdot \sin \omega t / \cos A). \quad [31]$$

To establish the correct quadrant, first calculate the hour angle (H_w) at which the sun is due west of the observer:

$$H_w = \cos^{-1}(\tan \delta \cdot \cotan \theta). \quad [32]$$

Then, if $|\omega t|$ is less than or equal to H_w , A' is in the correct quadrant. However, if $|\omega t|$ is greater than H_w , A' is placed in the correct quadrant by

$$A' = \text{sign}(180, \omega t) - A' \quad [33]$$

where $\text{sign}(180, \omega t)$ is 180° if ωt is greater than or equal to 0 and -180° if ωt is less than 0° .

If $|\theta|$ is less than or equal to $|\delta|$, three possible cases exist. In the first case, $|\omega t|$ is greater than 0° but less than 180° . If δ equals 0° , then θ equals 0° and

$$A' = \text{sign}(90, \omega t). \quad [34]$$

If δ is not equal to 0° , then A' is calculated as in equation [31]. A' from equation [31] is in the correct quadrant if δ is less than 0° . For δ greater than 0° , A' is determined using equation [33].

In the second case, ωt equals 0° . If $|\theta|$ equals $|\delta|$, then when θ is greater than or equal to 0° , A' is 0° ; when θ is less than 0° , A' is 180° . If $|\theta|$ is less than $|\delta|$, then when δ is greater than 0° , A' is 180° ; when δ is less than 0° , A' is 0° .

In the third case, $|\omega t|$ equals 180° . If $|\theta|$ equals $|\delta|$, then when θ is greater than or equal to 0° , A' is 180° , and when θ is less than 0° , A' is 0° . If $|\theta|$ is less than $|\delta|$, then when δ is greater than 0° , A' is 180° , and when δ is less than 0° , A' is 0° .

To determine the effects of intervening terrain, the terrain altitude and azimuth must be determined for the opposite horizon (east horizon for a slope having an azimuth of 0° to 180° and west horizon for a slope having an azimuth of 180° to 360°). This may be done by mapping the elevation angle of the horizon at suitably spaced azimuths or by calculating the elevation angle using topography maps.

At some sites, such as lower, south-facing slopes in valleys (northern hemisphere), the entire horizon between the points of sunrise and sunset must be mapped to determine if the sun is obstructed by the opposite ridge. The portion of the obstructed horizon which must be mapped and the azimuth difference between points at which elevation angle is measured will depend upon the particular site in question and the accuracy required.

After the terrain profile is determined, solar altitude is compared with the terrain elevation at various solar azimuths to determine the times at which the obstruction occludes the sun. For instantaneous irradiance calculations, M_i or R_i is set to zero if the solar altitude is less than the elevation angle of the terrain. For total daily irradiance, M_t or R_t are calculated using t'_1 and t'_2 adjusted for the effects of slope at the given site and for the effects of terrain elevation.

Explanation of Irradiance Figures

A series of graphs was prepared to illustrate how the instantaneous irradiance multiplier is influenced by latitude, date, azimuth (aspect), inclination (percent slope), and time of day. Three figures are presented for each of three latitudes—30°, 40°, and 50° N. These figures are for June 22, March or September 22, and December 22, representing the summer solstice, vernal and autumnal equinoxes, and winter solstice in the northern hemisphere. Slopes in degrees are as follows: 0% (0°), 20% (11.3°), 50% (26.6°), 100% (45°), 200% (63.4°), 400% (76.0°). Effects of obstructed opposite horizons are not taken into account.

The figures illustrate the complex geometric relationships among latitude, solar declination, azimuth, inclination, and hour angle. Total day length decreases from June 22 to December 22, but the change in day length is greatest at high latitudes. For example, at 30° N latitude, day length decreases from 13.9 hours to 10.1 hours (figs. 2 and 4), while at 50° N latitude, day length decreases from 16.2 hours to 7.8 hours (figs. 8 and 10).

Vertical dashed lines in the figures represent sunrises or sunsets at the true horizon for surfaces tilted east or west. For east-facing surfaces (azimuth of 90°), an increase in slope increases the intensity of radiation immediately after the sun rises past the horizon. However, steeper slopes also experience earlier sunsets because the effective horizon is elevated. Opposite effects occur on west-facing slopes.

Instantaneous irradiance of north-facing slopes (azimuth of 0°) at solar noon decreases as the inclination increases. At steep inclinations, the sun may set in the morning and rise in the afternoon during summer months; this effect becomes more pronounced at higher latitudes (compare figs. 2, 5, 8). During winter months, low sun angles prevent any direct beam radiation from reaching steep, north-facing slopes (figs. 4, 7, 10). On south-facing slopes (azimuth of 180°), instantaneous irradiance is often higher than on horizontal surfaces,

because the normal to the slope is close to the solar beam pathway.

Figures 2 through 10 are presented only to demonstrate the effects of various factors which influence instantaneous irradiance. Through proper use of the equations given above, the user may calculate instantaneous irradiance for any latitude, date, azimuth, or inclination. In the southern hemisphere, instantaneous irradiance multipliers for slopes facing directly east or west are similar to those shown in the figures. However, for north- and south-facing slopes, the graphs are reversed.

The total irradiance multiplier (M_t), calculated with appropriate equations above, is represented by the area under the curves between sunrise and sunset. Frank and Lee (1966) tabulated total daily irradiance for latitudes between 30° and 50° N, 16 azimuths, and slopes from 0% to 100% for various times of the year. Their values were calculated with a solar irradiance constant of $2.0 \text{ cal} \cdot \text{cm}^{-2} \cdot \text{min}^{-1}$. Irradiance multipliers for total solar irradiance (M_t) are equal to Frank and Lee's values divided by 120 (from $2 \text{ cal} \cdot \text{cm}^{-2} \cdot \text{min}^{-1}$ times 60 min). Tabulated values of Buffo et al. (1972) include an atmospheric transmission coefficient and cannot be compared directly with those given here or in Frank and Lee (1966).

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Latitude: 30 deg N

Date: June 22

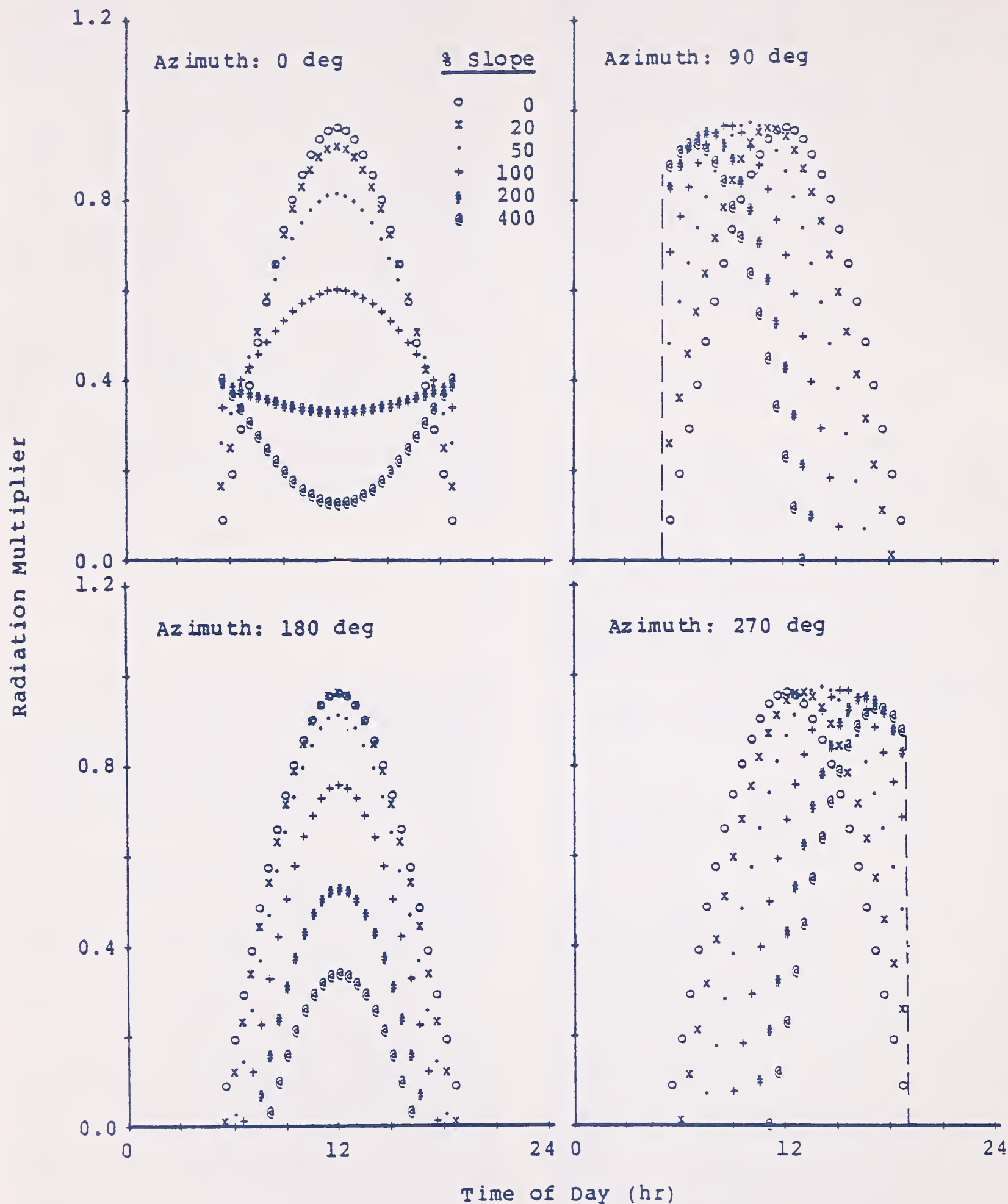


Figure 2.—Instantaneous irradiance multiplier for 30° N latitude on June 22.

Latitude: 30 deg N

Date: March 22 or September 22

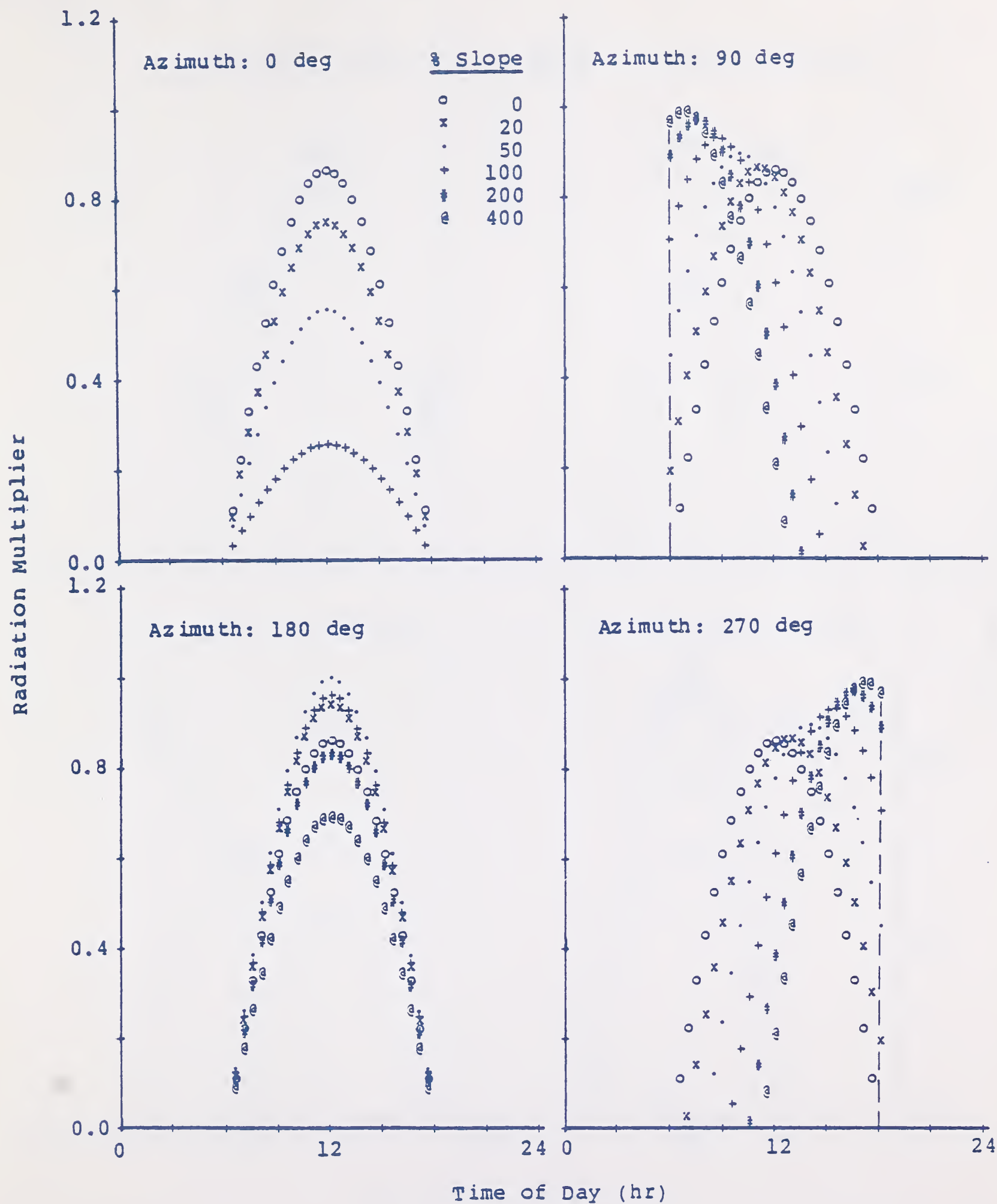


Figure 3.—Instantaneous irradiance multiplier for 30° N latitude on March 22 or September 22.

Latitude: 30 deg N Date: December 22

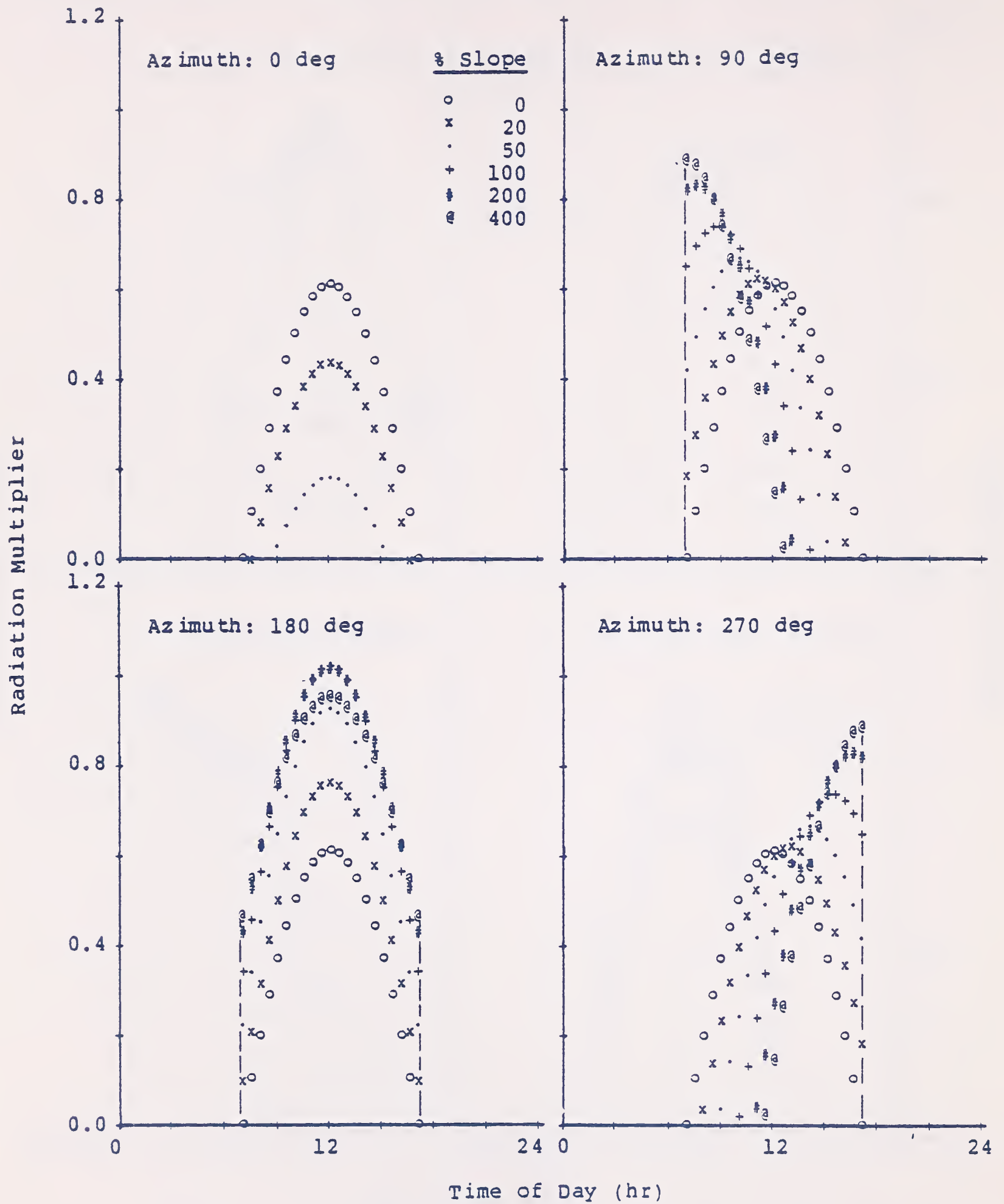


Figure 4.—Instantaneous irradiance multiplier for 30° N latitude on December 22.

Latitude: 40 deg N

Date: June 22

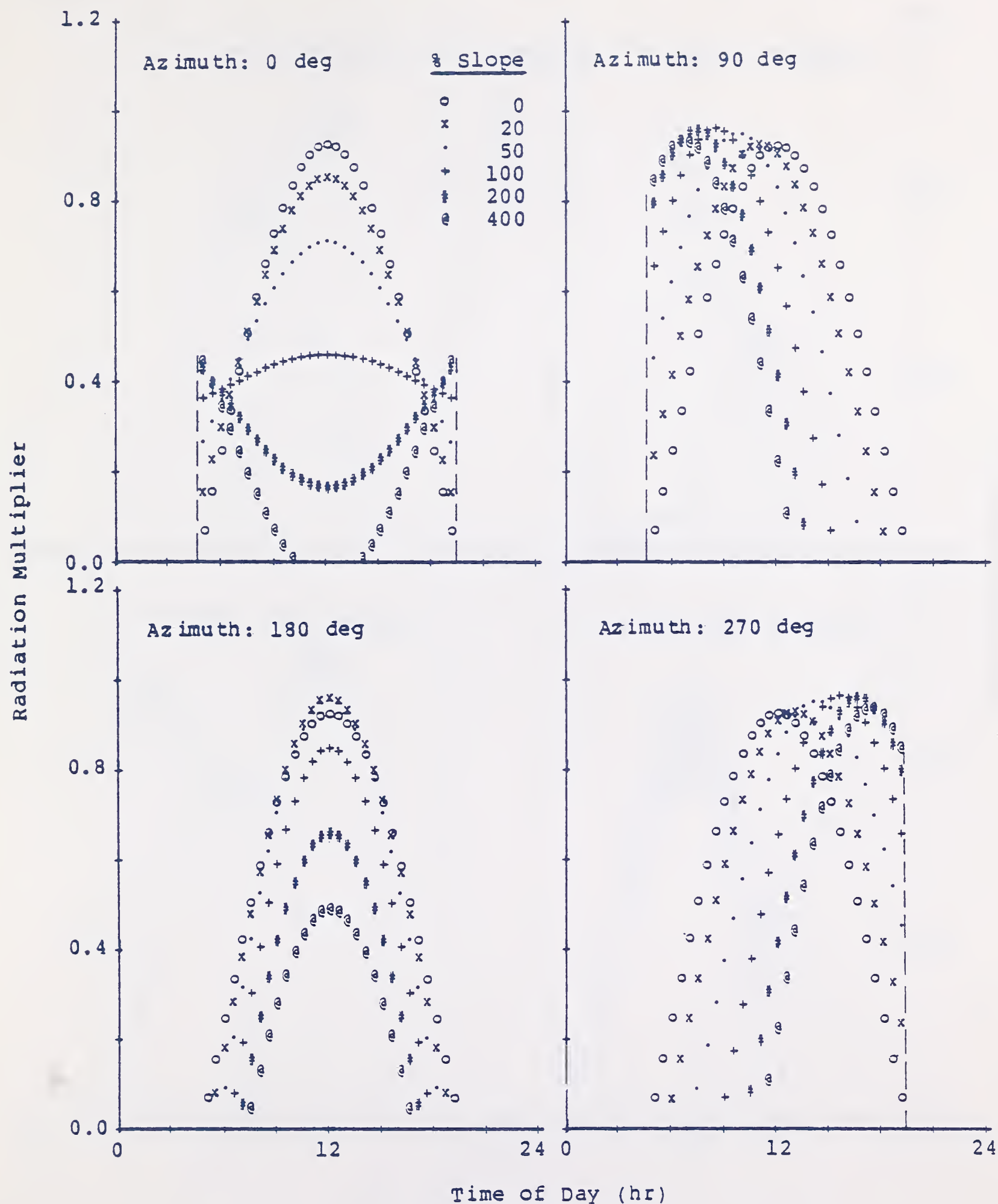


Figure 5.—Instantaneous irradiance multiplier for 40° N latitude on June 22.

Latitude: 40 deg N

Date: March 22 or September 22

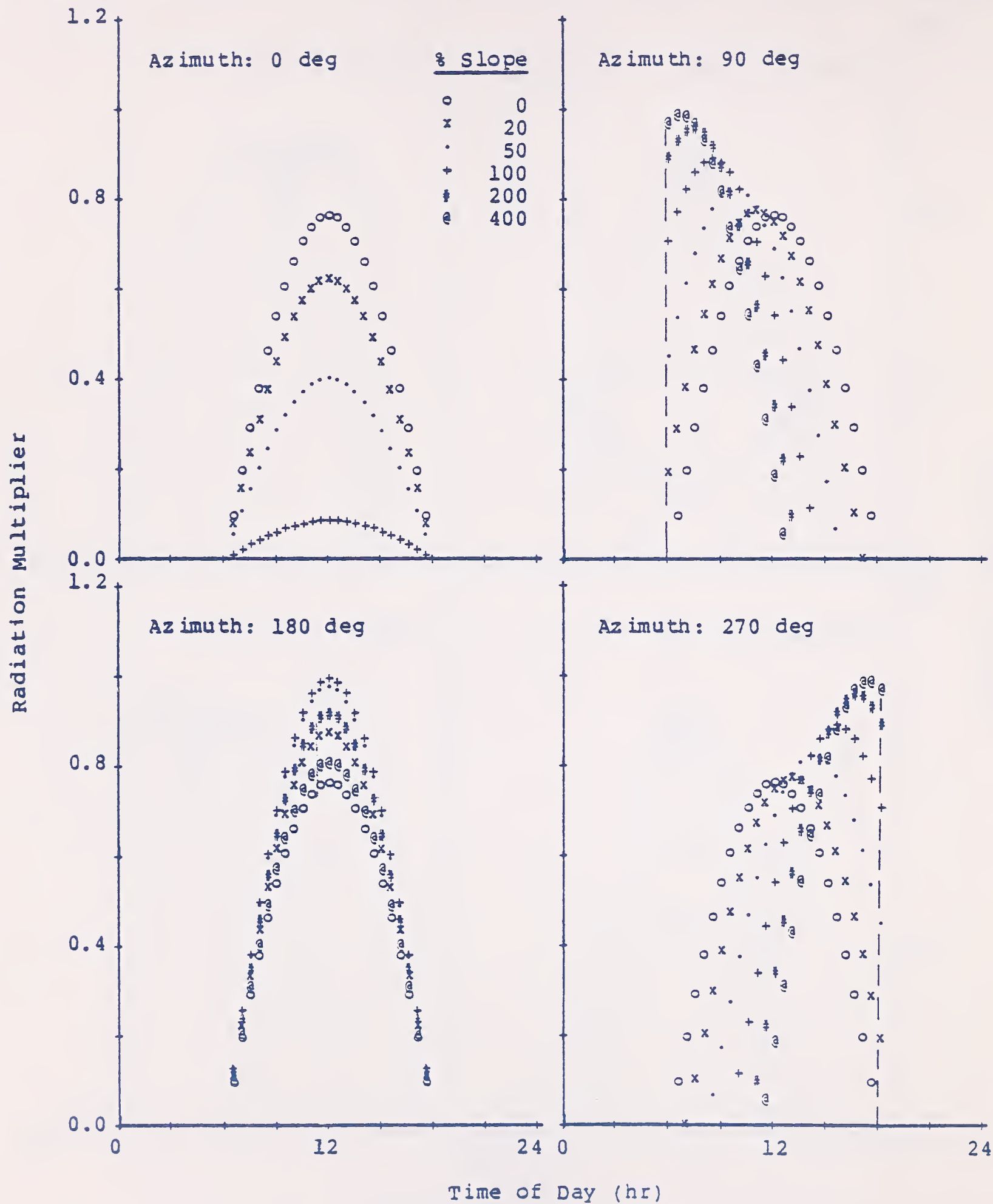


Figure 6.—Instantaneous irradiance multiplier for 40° N latitude on March 22 and September 22.

Latitude: 40 deg N

Date: December 22

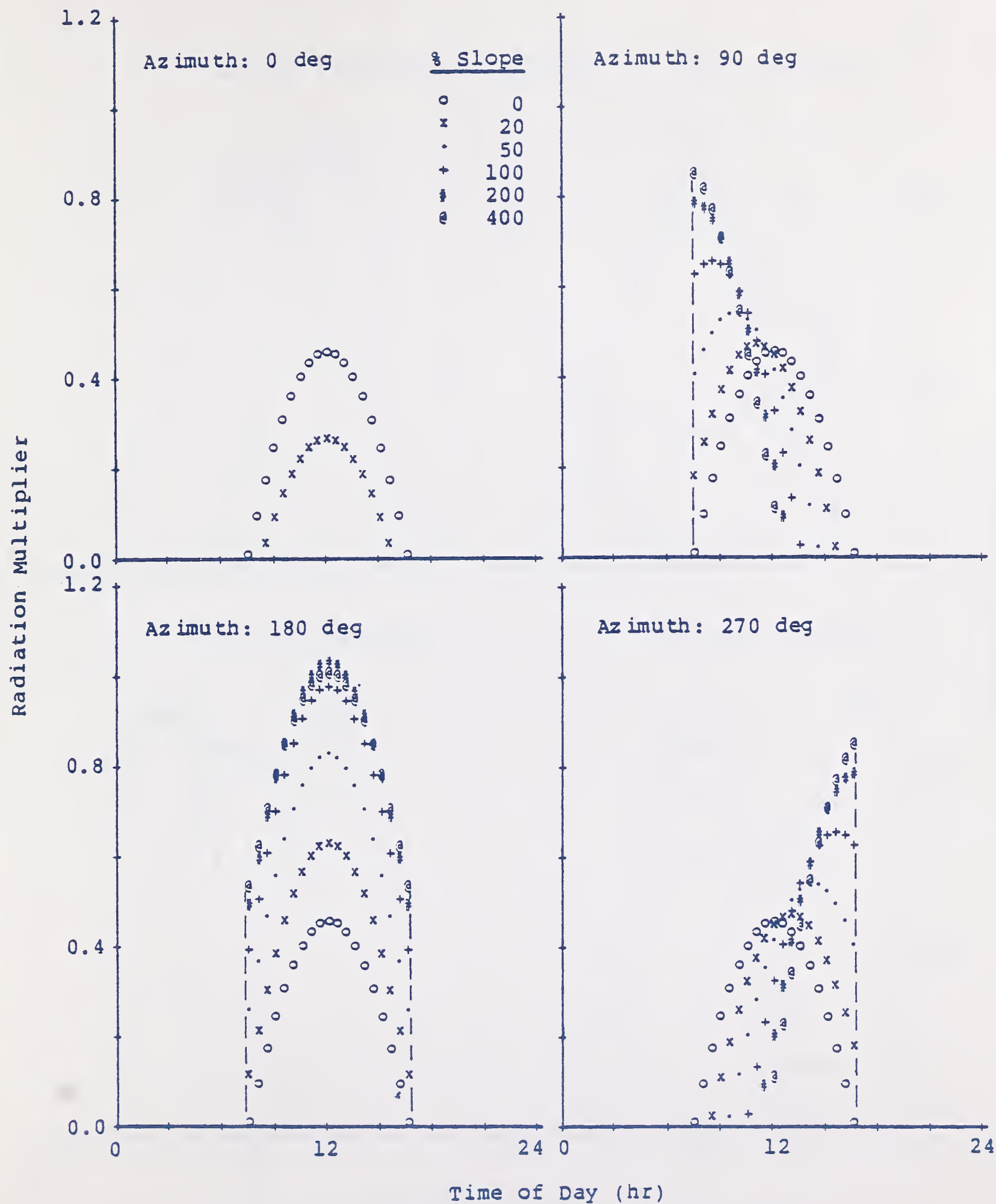


Figure 7.—Instantaneous irradiance multiplier for 40° N latitude on December 22.

Latitude: 50 deg N

Date: June 22

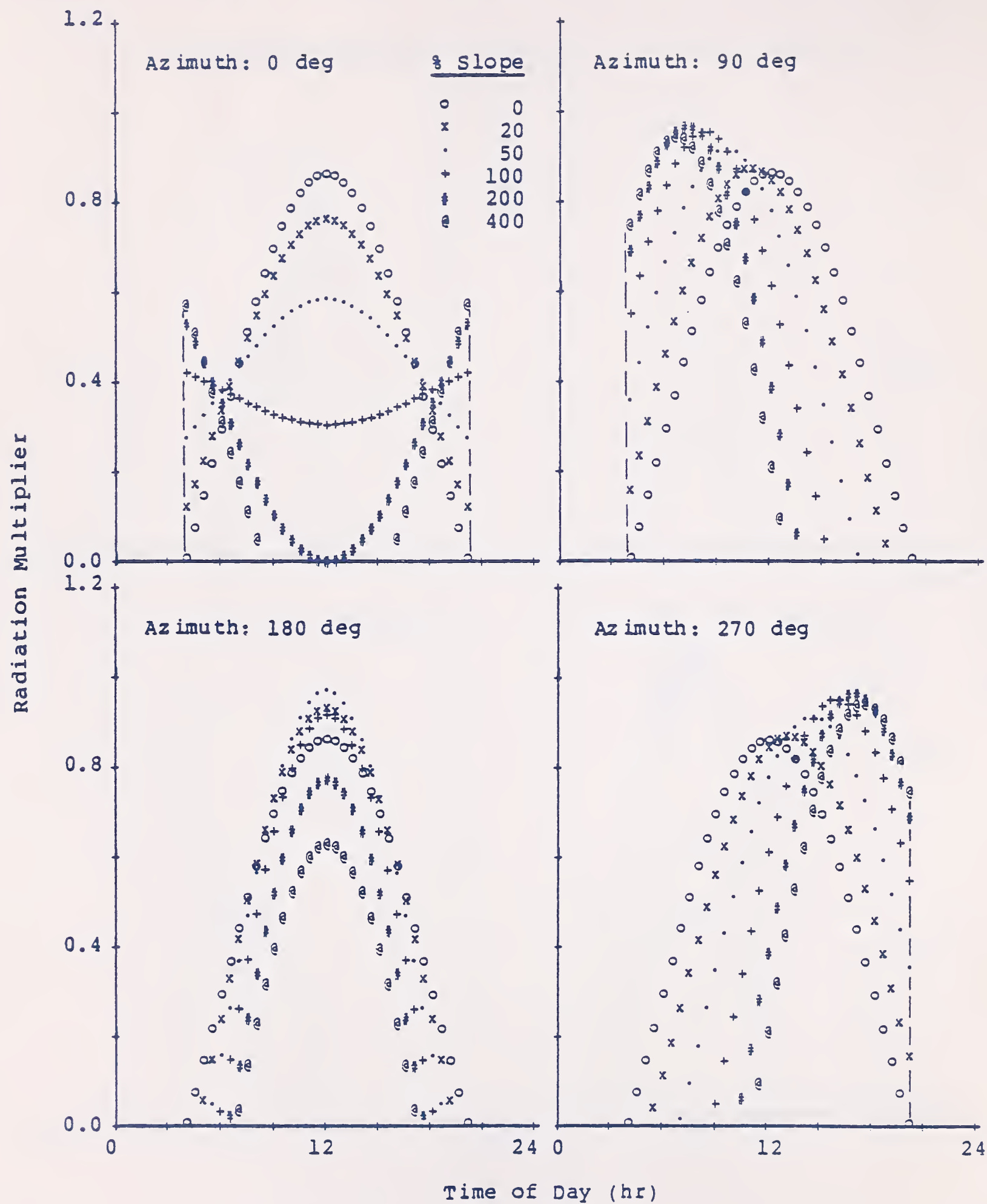


Figure 8.—Instantaneous irradiance multiplier for 50° N latitude on June 22.

Latitude: 50 deg N

Date: March 22 or September 22

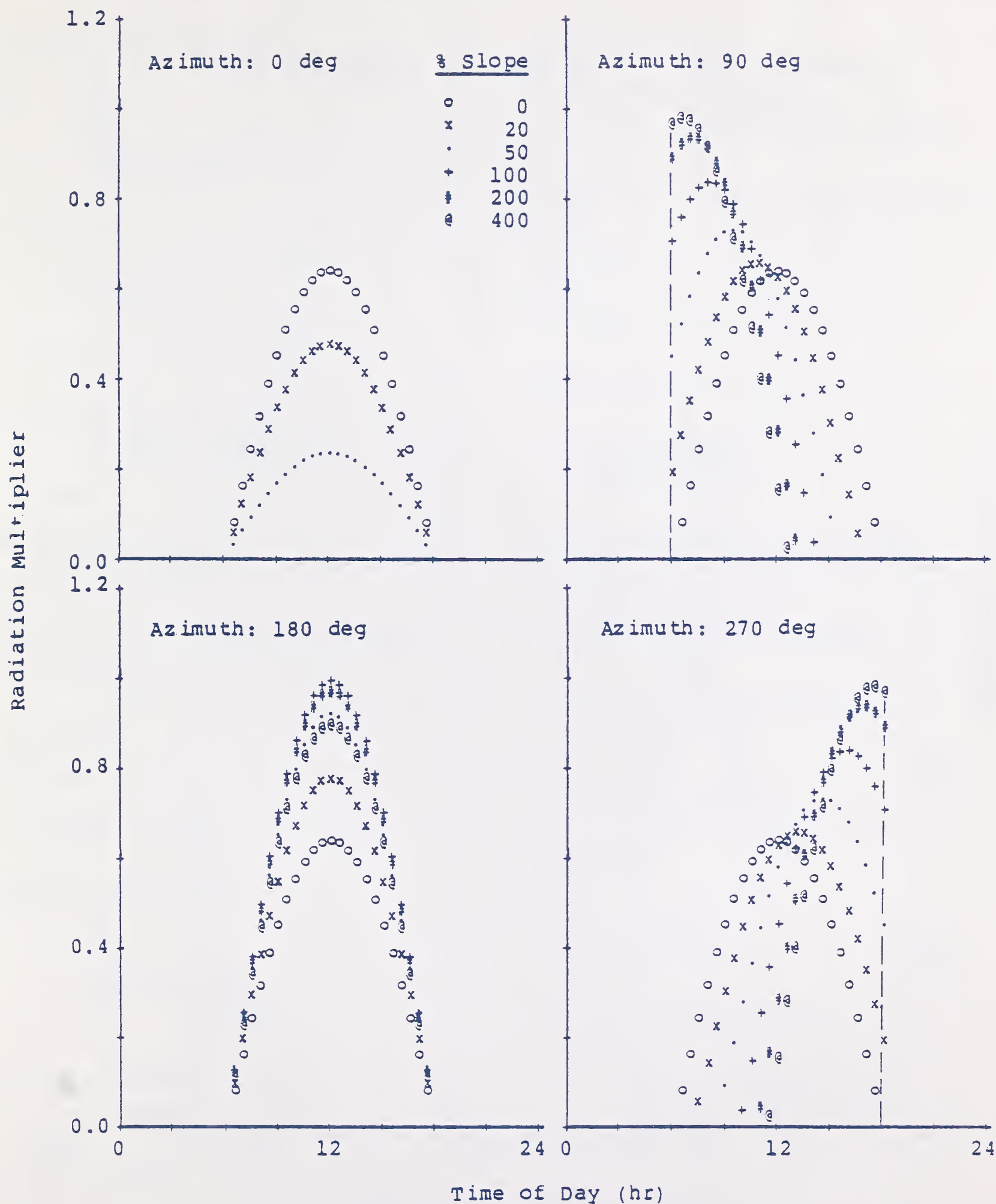


Figure 9.—Instantaneous irradiance multiplier for 50° N latitude on March 22 or September 22.

Latitude: 50 deg N

Date: December 22

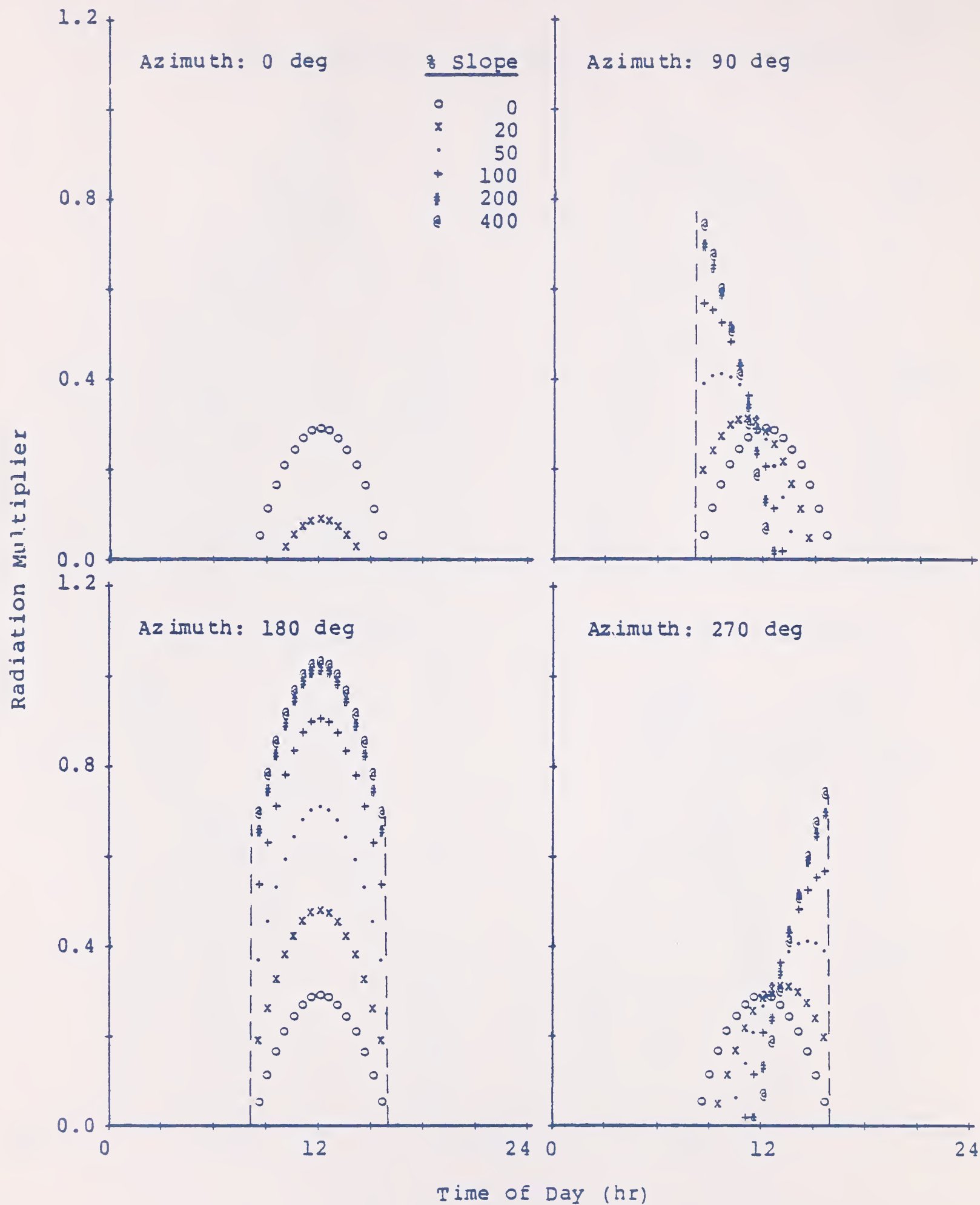


Figure 10.—Instantaneous irradiance multiplier for 50° N latitude on December 22.

Appendix A

Example 1.—Horizontal Surface

Date	June 22
θ	40°
δ	23.5°
R_o	$1360 \text{ W} \cdot \text{m}^{-2} (1.95 \text{ cal} \cdot \text{cm}^{-2} \cdot \text{min}^{-1})$
t	$-2.00 \text{ hr (1000 hr)}$

Times of Sunrise and Sunset:

$$\begin{aligned}\omega t &= \cos^{-1}(-\tan 40^\circ \cdot \tan 23.5^\circ) \\ t_2 &= 7.43 \text{ hr (1926 hr)} \\ t_1 &= -7.43 \text{ hr (0434 hr)}\end{aligned}$$

Instantaneous Irradiance Multiplier:

$$\begin{aligned}\omega t &= -2 \cdot 15^\circ = -30^\circ \\ M_i &= 1.0[\sin 40^\circ \cdot \sin 23.5^\circ \\ &\quad + \cos 40^\circ \cdot \cos 23.5^\circ \cdot \cos(-30^\circ)] \\ &\quad \div (0.999847 + 0.001406 \cdot 23.5^\circ) \\ M_i &= 0.8372\end{aligned}$$

Instantaneous Irradiance:

$$\begin{aligned}R_i &= 0.8372 \cdot 1360 \text{ W} \cdot \text{m}^{-2} \\ R_i &= 1138.6 \text{ W} \cdot \text{m}^{-2}\end{aligned}$$

Total Daily Irradiance Multiplier:

$$\begin{aligned}M_t &= 1.0[2 \cdot 7.43 \cdot \sin 40^\circ \cdot \sin 23.5^\circ \\ &\quad + (12/\pi) \cdot \cos 40^\circ \cdot \cos 23.5^\circ \cdot 2 \cdot \sin 111.4^\circ] \\ &\quad \div 1.0329 \\ M_t &= 8.5251\end{aligned}$$

Total Daily Irradiance:

$$\begin{aligned}R_t &= 8.5251 \cdot 60 \cdot 1360 \\ R_t &= 6.956 \times 10^5 \text{ W} \cdot \text{m}^{-2}\end{aligned}$$

Example 2.—Tilted Surface

Date	September 22
θ	40°
δ	0°
R_o	$1360 \text{ W} \cdot \text{m}^{-2}$
t	$1.00 \text{ hr (1300 hr)}$
a	90°
i	40%

Inclination in Degrees:

$$\begin{aligned}i &= \tan^{-1}(40/100) \\ i &= 21.8^\circ\end{aligned}$$

Equivalent Latitude:

$$\begin{aligned}\theta' &= \sin^{-1}(\sin 21.8^\circ \cdot \cos 90^\circ \cdot \cos 40^\circ \\ &\quad + \cos 21.8^\circ \cdot \sin 40^\circ) \\ \theta' &= 36.64^\circ\end{aligned}$$

Hour Angle Correction:

$$\begin{aligned}\alpha &= \tan^{-1}[(\sin 90^\circ \cdot \sin 21.8^\circ) \\ &\quad \div (\cos 21.8^\circ \cdot \cos 40^\circ \\ &\quad - \cos 90^\circ \cdot \sin 21.8^\circ \cdot \sin 40^\circ)] \\ \alpha &= 27.57^\circ\end{aligned}$$

Sunrise and Sunset Hour Angles on Equivalent Slope:

$$\begin{aligned}\omega t' &= \cos^{-1}(-\tan 36.64^\circ \cdot \tan 0^\circ) \\ \omega t' &= 90^\circ \\ \text{Also,} \\ \omega t &= \cos^{-1}(-\tan 40^\circ \cdot \tan 0^\circ) \\ \omega t &= 90^\circ\end{aligned}$$

Sunrise and Sunset for the Given Slope:

$$\begin{aligned}t_1 &= [\max(-90^\circ, -117.57^\circ)] / 15 \\ t_1 &= -6.00 \text{ hr (0600 hr)} \\ t_2 &= [\min(90^\circ, 62.43^\circ)] / 15 \\ t_2 &= 4.16 \text{ hr (1610 hr)}\end{aligned}$$

Sunrise and Sunset for the Equivalent Slope:

$$\begin{aligned}t_1' &= [\max(-62.43^\circ, -90^\circ)] / 15 \\ t_1' &= -4.16 \text{ hr (0750 hr)} \\ t_2' &= [\min(117.57^\circ, 90^\circ)] / 15 \\ t_2' &= 6.00 \text{ hr (1800 hr)}\end{aligned}$$

Instantaneous Irradiance Multiplier:

$$\begin{aligned}\omega t'' &= 15 \cdot 1.00 + 27.57^\circ \\ \omega t'' &= 42.57^\circ \\ M_i &= 1.0[(\sin 36.64^\circ \cdot \sin 0^\circ) \\ &\quad + (\cos 36.64^\circ \cdot \cos 0^\circ \cdot \cos 42.57^\circ)] \\ &\quad \div (0.999847 + 0.001406 \cdot 0^\circ) \\ M_i &= 0.5910\end{aligned}$$

Instantaneous Irradiance:

$$\begin{aligned}R_i &= 0.5910 \cdot 1360 \text{ W} \cdot \text{m}^{-2} \\ R_i &= 803.8 \text{ W} \cdot \text{m}^{-2}\end{aligned}$$

Total Daily Irradiance Multiplier:

$$\begin{aligned}M_t &= 1.0\{(6.00 + 4.16) \cdot \sin 36.64^\circ \cdot \sin 0^\circ \\ &\quad + (12/\pi) \cdot \cos 36.64^\circ \cdot \cos 0^\circ \\ &\quad \cdot [\sin 90^\circ - \sin(-62.43^\circ)]\} \\ &\quad \div (0.999847 + 0.001406 \cdot 0^\circ) \\ M_t &= 5.7827\end{aligned}$$

Total Daily Irradiance:

$$\begin{aligned}R_t &= 5.7827 \cdot 60 \cdot 1360 \text{ W} \cdot \text{m}^{-2} \\ R_t &= 4.719 \times 10^5 \text{ W} \cdot \text{m}^{-2}\end{aligned}$$

Example 3.—Double Sunrise and Sunset

Date	June 22
θ	50°
δ	23.5°
a	0°
i	250%

Inclination in Degrees:

$$i = \tan^{-1}(250/100)$$
$$i = 68.2^\circ$$

Equivalent Latitude:

$$\theta' = \sin^{-1}(\sin 68.2^\circ \cdot \cos 0^\circ \cdot \cos 50^\circ + \cos 68.2^\circ \cdot \sin 50^\circ)$$

$$\theta' = 61.8^\circ \text{ (note: this latitude is on the opposite side of the North Pole from the given slope; see calculation for } \alpha \text{ below)}$$

Hour Angle Correction:

$$\theta + i = 50 + 68.2$$

$$\theta + i = 118.2$$

$$\text{Because } (\theta + i) > 90^\circ, \alpha = 180^\circ$$

$$\omega t = \cos^{-1}(-\tan 50^\circ \cdot \tan 23.5^\circ)$$

$$\omega t = 121.21^\circ$$

$$\omega t' = \cos^{-1}(-\tan 61.8^\circ \cdot \tan 23.5^\circ)$$

$$\omega t' = 144.19^\circ$$

Times of Double Sunrise and Sunset at the Given Slope:

$$\text{first } t_1 = -121.21^\circ/15$$

$$\text{first } t_1 = -8.08 \text{ hr (0355 hr)}$$

$$\text{first } t_2 = (144.19^\circ - 180^\circ)/15$$

$$\text{first } t_2 = -2.39 \text{ hr (0937 hr)}$$

$$\text{second } t_1 = (360^\circ - 144.19^\circ - 180^\circ)/15$$

$$\text{second } t_1 = 2.39 \text{ hr (1423 hr)}$$

$$\text{second } t_2 = 121.21^\circ/15$$

$$\text{second } t_2 = 8.08 \text{ hr (2005 hr)}$$

Times of Double Sunrise and Sunset at the Equivalent Slope:

$$\text{first } t'_1 = -144.19^\circ/15$$

$$\text{first } t'_1 = -9.61 \text{ hr (0223 hr)}$$

$$\text{first } t'_2 = (121.21^\circ + 180^\circ - 360^\circ)/15$$

$$\text{first } t'_2 = -3.92 \text{ hr (0805 hr)}$$

$$\text{second } t'_1 = (-121.21^\circ + 180^\circ)/15$$

$$\text{second } t'_1 = 3.92 \text{ hr (1555 hr)}$$

$$\text{second } t'_2 = 144.19^\circ/15$$

$$\text{second } t'_2 = 9.61 \text{ hr (2137 hr)}$$

Appendix B

THIS PROGRAM ILLUSTRATES THE USAGE OF SUBPROGRAMS

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JULDAT
DEC
RADVEC
RAD
GT
GI
SOLALT
SOLAZ

BY CALCULATING A TOTAL IRRADIATION MULTIPLIER
AND A TABLE OF INSTANTANEOUS IRRADIATION
MULTIPLIERS, SOLAR ALTITUDES AND SOLAR AZIMUTHS
AT 1/4 HOUR INTERVALS FOR 24 HOURS BEGINNING
AT SOLAR MIDNIGHT (12 HOURS PAST SOLAR NOON).

REQUIRED INPUT:

LATITUDE OF THE SITE IN DEGREES
AZIMUTH OF THE SITE MEASURED IN DEGREES
CLOCKWISE FROM NORTH
SLOPE OF THE SITE IN %
INTEGER VALUE OF THE MONTH OF THE PREDICTION
INTEGER VALUE OF THE DAY OF MONTH OF THE
PREDICTION

ALL OUTPUT TIMES ARE IN HOURS BEFORE (LT 0) OR AFTER (GT 0)
SOLAR NOON. ALL OUTPUT HOUR ANGLES ARE IN DEGREES
CLOCKWISE FROM THE UPPER MERIDIAN. SOLAR AZIMUTHS ARE GIVEN
IN DEGREES CLOCKWISE FROM SOUTH.

COMMON/PIE/PI,PI2,PI12,PI180
DIMENSION GTIME(4),ETIME(4)
INTEGER DAY
REAL LAT
LOGICAL SHADED,DOUBLE
REAL IRRAD

PI=4.*ATAN(1.)
PI2=PI/2.
PI12=PI/12.
PI180=PI/180.

READ AND PRINT USER INPUT

10 PRINT 100
10J FORMAT(1H0,'ENTER LATITUDE IN DEGREES')
READ(5,*,END=250) LAT
PRINT 110

SET VARIOUS MULTIPLIES OF PI FOR
COMMON BLOCK PIE

55 110 FCRMAT(1H0,'ENTER AZIMUTH AS DEGREES CLOCKWISE FROM NORTH')
56 READ(5,*) AZMUTH
57 PRINT 120
58 120 FORMAT(1H0,'ENTER % SLOPE')
59 READ(5,*) SLOPE
60 PRINT 130
61 130 FORMAT(1H0,'ENTER MONTH OF PREDICTION')
62 READ(5,*) MONTH
63 PRINT 140
64 140 FORMAT(1H0,'ENTER DAY OF PREDICTION')
65 READ(5,*) DAY
66 C
67 PRINT 150, LAT,AZMUTH,SLOPE,MONTH,DAY
68 150 FORMAT(1H0/1H ,%LATITUDE - %,F6.1,% DEGREES%/1H ,
69 & %AZIMUTH - %,F6.1,% DEGREES%/1H ,
70 & %SLOPE - %,F6.1,% %/1H ,
71 & %DATE OF PREDICTION - %,I2,%,/%,I2)
72 C
73 C
74 C
75 C
76 C
77 N=JULDAT(MONTH,DAY)
78 D=DEC(N)
79 R2=RADVEC(D)
80 C
81 C
82 PRINT 160, N,D/PI180,SQRT(R2)
83 160 FORMAT(1H0,'JULIAN DATE - %,I4/1H ,
84 & %DECLINATION - %,F6.1,% DEGREES%/1H ,
85 & %RADIUS VECTOR - %,F8.5)
86 C
87 C
88 C
89 C
90 RLAT=LAT*PI180
91 A=AZMUTH*PI180
92 SLP=ATAN(SLOPE/100.)
93 C
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107 170, THETA/PI180,ALPHA/PI180
108 170 FORMAT(1H0,'LATITUDE OF EQUIVALENT SLOPE - %,F6.1,% DEGREES%/1H ,
109 & %CHANGE IN LONGITUDE - %,F7.1,% DEGREES')
110 C
110 C
110 C

SET N - JULIAN DATE
D - SOLAR DECLINATION
R2- RADIUS VECTOR SQUARED

DIVISION BY PI180 CONVERTS
RADIANS TO DEGREES

CONVERT LATITUDE,AZIMUTH, AND
SLOPE TO RADIAN MEASURE

CALL TO RAD ROUTINE TO CALCULATE
THETA - EQUIVALENT LATITUDE
ALPHA - CHANGE IN LONGITUDE
GTIME - SUNRISE SUNSET AT GIVEN
SLOPE
ETIME - SUNRISE SUNSET AT THE
EQUIVALENT SLOPE
AND DETERMINE THE NUMBER OF
SUNRISE SUNSET PERIODS

CALL RAD(RLAT,A,SLP,D,THETA,ALPHA,SHADED,DOUBLE,GTIME,ETIME)

CALCULATE RADIATION MULTIPLIERS


```

111 C      AND GENERATE TABLE
112 C
113 C      IF(SHADED) THEN
114 C
115 C          GO TO NEXT PROBLEM SINCE
116 C          SUN DOES NOT RISE
117 C
118 C          PRINT 180
119 C          FORMAT(1H0/1H ,*SLOPE IS SHADED FOR ENTIRE DAY*)
120 C
121 C      ELSE
122 C
123 C          AT LEAST 1 SUNRISE SUNSET PERIOD
124 C
125 C          ONE SUNRISE SUNSET PERIOD
126 C
127 C          DIVISION BY PI12 CONVERTS
128 C          HOUR ANGLES IN RADIANS TO
129 C          HOURS BEFORE OR AFTER SOLAR
130 C          NOON
131 C
132 C          PRINT 190, (GTIME(1)/PI12,I=1,2),(ETIME(1)/PI12,I=1,2)
133 C          FORMAT(1H0/1H ,*GIVEN SLOPE SUNRISE*,F6.1/1H ,
134 C          12X,
135 C          *SUNSET*,F6.1/1H0,
136 C          *EQUIVALENT SLOPE SUNRISE*,F6.1/1H ,
137 C          17X,
138 C          *SUNSET*,F6.1)
139 C
140 C      ELSE
141 C
142 C          DOUBLE DAY - IE TWO SUNRISE
143 C          SUNSET PERIODS
144 C
145 C          PRINT 200, (GTIME(1)/PI12,I=1,4),(ETIME(1)/PI12,I=1,4)
146 C          FORMAT(1H0/1H ,*DOUBLE SUNRISE SUNSET*/1H0,
147 C          *GIVEN SLOPE FIRST SUNRISE*,F6.1/1H ,
148 C          12X,
149 C          *FIRST SUNSET*,F6.1/1H ,
150 C          12X,
151 C          *SECOND SUNRISE*,F6.1/1H ,
152 C          12X,
153 C          *SECOND SUNSET*,F6.1/1H0,
154 C          *EQUIVALENT SLOPE FIRST SUNRISE*,F6.1/1H ,
155 C          17X,
156 C          *FIRST SUNSET*,F6.1/1H ,
157 C          12X,
158 C          *SECOND SUNRISE*,F6.1/1H ,
159 C          17X,
160 C          *SECOND SUNSET*,F6.1/1H ,
161 C          17X,
162 C          *SECOND SUNSET*,F6.1)
163 C
164 C      END IF
165 C
166 C          SET TRAD -TOTAL DAILY IRRADIATION
167 C          MULTIPLIER FOR FIRST
168 C          SUNRISE-SET PERIOD
169 C
170 C          TRAD=QT(ETIME(1),ETIME(2),THETA,D,R2)
171 C
172 C          IF(DOUBLE)   TRAD=TRAD+QT(ETIME(3),ETIME(4),THETA,D,R2)
173 C
174 C          PRINT 210, TRAD
175 C          FORMAT(1H0/1H ,*MULTIPLIER FOR TOTAL IRRADIATION - *,F6.2)
176 C          PRINT 220
177 C          FORMAT(1H0/1H ,36X,*INSTANTANEOUS*/1H ,
178 C          10X,*HOUR*,5X,*SOLAR*,4X,*SOLAR*,4X,
179 C          *IRRADIATION*/1H ,
180 C          *TIME*,5X,*ANGLE*,3X,*ALTITUDE*,2X,
181 C          *AZIMUTH*,4X,*MULTIPLIER*/)
182 C
183 C          CALCULATE TABLE OF SOLAR
184 C          ALTITUDES AND AZIMUTHS FOR
185 C          15 MINUTE INTERVALS
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THIS ROUTINE WILL CALCULATE

THP - LATITUDE OF THE EQUIVALENT SLOPE IN RADIAN'S
      (THP LT 0 FOR SOUTHERN HEMISPHERE THP GT 0 FOR
      NORTHERN HEMISPHERE)
ALPHA - CHANGE IN LONGITUDE FROM THE GIVEN SLOPE TO THE
      EQUIVALENT SLOPE IN RADIAN'S (-180 LT ALPHA LE 180)
SHADED - IF SHADED=.TRUE., THE GIVEN SLOPE IS SHADED
DOUBLE - IF DOUBLE=.TRUE., THE GIVEN SLOPE EXPERIENCES
      A DOUBLE SUNRISE-SUNSET. IF DOUBLE=.FALSE., AND
      SHADED=.FALSE., THE GIVEN SLOPE EXPERIENCES ONE
      SUNRISE AND ONE SUNSET
GTIME(1) - HOUR ANGLE OF THE FIRST SUNRISE AT THE
      GIVEN SLOPE IN RADIAN'S (-180 LE GTIME(1) LE 180)
GTIME(2) - HOUR ANGLE OF THE FIRST SUNSET AT THE
      GIVEN SLOPE IN RADIAN'S (-180 LE GTIME(2) LE 180)
GTIME(3) - HOUR ANGLE OF THE SECOND SUNRISE AT THE
      GIVEN SLOPE IN RADIAN'S (-180 LE GTIME(3) LE 180)
GTIME(4) - HOUR ANGLE OF THE SECOND SUNSET AT THE
      GIVEN SLOPE IN RADIAN'S (-180 LE GTIME(4) LE 180)
ETIME(1) - HOUR ANGLE OF THE FIRST SUNRISE AT THE
      EQUIVALENT SLOPE IN RADIAN'S (-180 LE ETIME(1) LE 180)
ETIME(2) - HOUR ANGLE OF THE FIRST SUNSET AT THE
      EQUIVALENT SLOPE IN RADIAN'S (-180 LE ETIME(2) LE 180)
ETIME(3) - HOUR ANGLE OF THE SECOND SUNRISE AT THE
      EQUIVALENT SLOPE IN RADIAN'S (-180 LE ETIME(3) LE 180)
ETIME(4) - HOUR ANGLE OF THE SECOND SUNSET AT THE
      EQUIVALENT SLOPE IN RADIAN'S (-180 LE ETIME(4) LE 180)

NOTE THAT IF SHADED=.TRUE., GTIME AND ETIME ARE MEANINGLESS.
IF SHADED=.FALSE., AND .DOUBLE=.FALSE., THEN GTIME(3), GTIME(4),
ETIME(3), AND ETIME(4) ARE MEANINGLESS.

ONCE THESE VARIABLES ARE COMPUTED THE USER MAY INCORPORATE
SUBPROGRAMS GI AND GT TO CALCULATE IRRADIATION MULTIPLIERS.

REQUIRED SUBPROGRAMS : NONE

PROGRAMMING NOTES :

USAGE - CALL RAD(TH,A,SLP,D,THP,ALPHA,SHADED,DOUBLE,GTIME,ETIME)

PARAMETER TYPES :

PARAMETERS TH,A,SLP,D,THP,ALPHA ARE SCALARS OF TYPE REAL
PARAMETERS SHADED,DOUBLE ARE SCALARS OF TYPE LOGICAL
PARAMETERS GTIME,ETIME ARE ONE DIMENSIONAL ARRAYS OF
      TYPE REAL AND DIMENSION 4
ALL VARIABLES ARE SINGLE PRECISION

COMMON BLOCKS - COMMON/PIE/PI,PI2,PI12,PI180 WHERE
      PI=4*ATAN(1)
      PI2=PI/2
      PI12=PI/12

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69 C      PI180=PI/180
70 C
71 C
72 C      SUBROUTINE RAD(TH,A,SLP,D,THP,ALPHA,SHADED,DOUBLE,GTIME,ETIME)
73 C      COMMON/PIE/PI,PI2,PI12,PI180
74 C      REAL TH,A,SLP,THP,ALPHA,GTIME(4),ETIME(4)
75 C      LOGICAL SHADED,DOUBLE
76 C
77 C      CALCULATE LAT OF EQUIV SLOPE - THP
78 C      THP=SIN(SLP)*COS(A)*COS(TH) + COS(SLP)*SIN(TH)
79 C      THP=ASIN(THP)
80 C
81 C      CALCULATE CHANGE IN LONGITUDE - ALPHA
82 C
83 C      IF(A .GT. 0.) THEN
84 C          SINA=SIN(A)
85 C          ALPHA=COS(SLP)*COS(TH) - COS(A)*SIN(SLP)*SIN(TH)
86 C          ALPHA=(SINA*SIN(SLP))/ALPHA
87 C          ALPHA=ATAN(ALPHA)
88 C          IF(SINA*ALPHA .LT. 0.) THEN
89 C              ALPHA=SIGN(PI,SINA)+ALPHA
90 C          END IF
91 C      ELSE
92 C          IF((ABS(TH)+SLP) .GT. PI2) THEN
93 C              ALPHA=PI
94 C          ELSE
95 C              ALPHA=0.
96 C          END IF
97 C      END IF
98 C
99 C
100 C      CALCULATE WT - HOUR ANGLE OF SUNSET AT GIVEN SLOPE
101 C      WTP - HOUR ANGLE OF SUNSET AT EQUIVALENT SLOPE
102 C
103 C      IF(ABS(TH)+ABS(D) .GT. PI2) THEN
104 C          IF(TH*0 .LT. 0.) THEN
105 C              WT=0.
106 C          ELSE
107 C              WT=PI
108 C          END IF
109 C      ELSE
110 C          WT=ACOS(-TAN(TH)*TAN(D))
111 C      END IF
112 C      IF(ABS(THP)+ABS(D) .GT. PI2) THEN
113 C          IF(THP*0 .LT. 0.) THEN
114 C              WTP=0.
115 C          ELSE
116 C              WTP=PI
117 C          END IF
118 C      ELSE
119 C          WTP=ACOS(-TAN(THP)*TAN(D))
120 C      END IF
121 C
122 C      CALCULATE SUNRISE SUNSET HOUR ANGLES
123 C      T1=-WTP
124 C

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125 T2=(WT+ALPHA)-(2.*PI)
126 T3=-WT+ALPHA
127 T4=WTP
128
129 IF (T4 .LE. T3) THEN
130   SHADED=.TRUE.
131   DOUBLE=.FALSE.
132 ELSE
133   SHADED=.FALSE.
134   IF (T1 .LT. T2 .AND. T2 .LT. T3 .AND. T3 .LT. T4) THEN
135     DOUBLE=.TRUE.
136   GTIME(1)=-WT
137   GTIME(2)=WTP-ALPHA
138   GTIME(3)=(2.*PI) - (WTP+ALPHA)
139   GTIME(4)=WT
140   ETIME(1)=T1
141   ETIME(2)=T2
142   ETIME(3)=T3
143   ETIME(4)=T4
144 ELSE
145   DOUBLE=.FALSE.
146   GTIME(1)=AMAX1(-WT,-WTP-ALPHA)
147   GTIME(2)=AMIN1(WT,WTP-ALPHA)
148   ETIME(1)=AMAX1(-WT+ALPHA,-WTP)
149   ETIME(2)=AMIN1(WT+ALPHA,WTP)
150 END IF
151 END IF
152 RETURN
153 END

1 C REAL FUNCTION DEC(N)
2 C
3 C PURPOSE :
4 C   GIVEN N - THE JULIAN DATE
5 C   THIS ROUTINE WILL CALCULATE THE CORRESPONDING SOLAR DECLINATION
6 C   IN RADIAN (-23.45 LE DEC LE 23.45)
7 C
8 C REQUIRED SUBPROGRAMS : NONE
9 C
10 C PROGRAMMING NOTES :
11 C
12 C   USAGE - N=JULDAT(MONTH,DAY)
13 C
14 C   PARAMETER TYPES :
15 C
16 C   PARAMETERS MONTH, DAY ARE SCALARS OF TYPE INTEGER
17 C
18 C   COMMON BLOCKS : NONE
19 C
20 C PROGRAMMING NOTES :
21 C
22 C   USAGE - D=DEC(N)
23 C
24 C   PARAMETER TYPES :
25 C
26 C   PARAMETER N IS A SCALAR OF TYPE INTEGER
27 C
28 C   COMMON BLOCKS - COMMON/PIE/PI,PI2,PI12,PI180 WHERE
29 C   PI=4*ATAN(1)
30 C   PI2=PI/2
31 C   PI12=PI/12
32 C   PI180=PI/180
33 C
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1 C REAL FUNCTION QI(TIME,SR,SS,THETA,ALPHA,DEC,R2)
2 C
3 C PURPOSE :
4 C GIVEN TIME - HOUR ANGLE IN RADIAN FROM SOLAR NOON
5 C AT THE GIVEN SLOPE (-180 LE TIME LE 180)
6 C SR - HOUR ANGLE IN RADIAN OF SUNRISE AT THE
7 C GIVEN SLOPE (-180 LE SR LE 180)
8 C SS - HOUR ANGLE IN RADIAN OF SUNSET AT THE
9 C GIVEN SLOPE (-180 LE SS LE 180)
10 C THETA - LATITUDE OF THE EQUIVALENT SLOPE IN RADIAN
11 C (-90 LE THETA LE 90)
12 C ALPHA - CHANGE IN LONGITUDE FROM THE GIVEN SLOPE TO THE
13 C EQUIVALENT SLOPE IN RADIAN (-180 LE ALPHA LE 180)
14 C DEC - SOLAR DECLINATION IN RADIAN FOR A PARTICULAR DAY
15 C (-23.45 LE DEC LE 23.45)
16 C R2 - THE SQUARED RATIO OF THE EARTH-SUN DISTANCE
17 C TO ITS MEAN FOR A PARTICULAR DAY
18 C THIS ROUTINE WILL CALCULATE AN INSTANTANEOUS IRRADIATION
19 C MULTIPLIER CORRESPONDING TO HOUR ANGLE TIME.
20 C
21 C REQUIRED SUBPROGRAMS : NONE
22 C
23 C PROGRAMMING NOTES :
24 C
25 C USAGE - IRRAD=QI(TIME,SR,SS,THETA,ALPHA,DEC,R2)
26 C
27 C PARAMETER TYPES :
28 C
29 C PARAMETERS TIME,SR,SS,THETA,ALPHA,DEC,R2 ARE SCALARS OF TYPE REAL
30 C
31 C COMMON BLOCKS : NONE
32 C
33 C
34 REAL FUNCTION QI(TIME,SR,SS,THETA,ALPHA,DEC,R2)
35 REAL TIME,SR,SS,THETA,ALPHA,DEC,R2
36 C
37 IF(TIME .LT. SR .OR. TIME .GT. SS) THEN
38 QI=0.
39 ELSE
40 WTPP=TIME + ALPHA
41 TERM1=SIN(THETA)*SIN(DEC)
42 TERM2=COS(THETA)*COS(DEC)*COS(WTPP)
43 QI=(1.0/R2) + (TERM1 + TERM2)
44 C
45 C THE RANGE OF QI IS MATHEMATICALLY
46 C LIMITED TO 0 LE QI LE 1. DUE TO
47 C ROUNDOFF AND/OR CANCELLATION ERRORS
48 C THE ABOVE STATEMENT MAY ASSIGN A
49 C VALUE LESS THAN 0 OR GREATER THAN 1
50 C TO QI. THE FOLLOWING STATEMENTS
51 C ALLEVIATE THE PROBLEM.
52 IF(QI .LT. 0.) QI=0.
53 IF(QI .GT. 1.) QI=1.
54 END IF
55 RETURN
56 END

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1 C REAL FUNCTION QI(SR,SS,THETA,DEC,R2)
2 C
3 C PURPOSE :
4 C GIVEN SR - HOUR ANGLE IN RADIAN OF SUNRISE AT THE
5 C SLOPE (-180 LE SR LE 180)
6 C SS - HOUR ANGLE IN RADIAN OF SUNSET AT THE
7 C SLOPE (-180 LE SS LE 180)
8 C THETA - LATITUDE OF THE SLOPE IN RADIAN
9 C (-90 LE THETA LE 90)
10 C DEC - SOLAR DECLINATION IN RADIAN FOR A PARTICULAR DAY
11 C (-23.45 LE DEC LE 23.45)
12 C R2 - THE SQUARED RATIO OF THE EARTH-SUN DISTANCE
13 C TO ITS MEAN FOR A PARTICULAR DAY
14 C THIS ROUTINE WILL CALCULATE A TOTAL IRRADIATION MULTIPLIER
15 C CORRESPONDING TO HOUR ANGLES SR SS.
16 C
17 C REQUIRED SUBPROGRAMS : NONE
18 C
19 C PROGRAMMING NOTES :
20 C
21 C USAGE - TRAD=QI(SR,SS,THETA,DEC,R2)
22 C
23 C PARAMETER TYPES :
24 C
25 C PARAMETERS SR,SS,THETA,DEC,R2 ARE SCALARS OF TYPE REAL
26 C
27 C COMMON BLOCKS - COMMON/PIE/PI,PI2,PI12,PI180 WHERE
28 C PI=4*ATAN(1)
29 C PI2=PI/2
30 C PI12=PI/12
31 C PI180=PI/180
32 C
33 C
34 REAL FUNCTION QI(SR,SS,THETA,DEC,R2)
35 REAL SR,SS,THETA,DEC,R2
36 COMMON/PIE/PI,PI2,PI12,PI180
37 C
38 SSMSR=(SS-SR)*12./PI
39 TERM1=SSMSR*SIN(THETA)*SIN(DEC)
40 TERM2=COS(THETA)*COS(DEC)*(SIN(SS)-SIN(SR))
41 QI=(1.0/R2)*(TERM1 + 12./PI*TERM2)
42 C
43 C THE RANGE OF QI IS MATHEMATICALLY
44 C LIMITED TO :
45 C 0 LE QI LE (1/R2)*(SSMSR+24./PI)
46 C
47 C DUE TO ROUNDING AND/OR CANCELLATION
48 C ERRORS THE ABOVE STATEMENT MAY ASSIGN
49 C A VALUE TO QI OUT OF THIS RANGE. THE
50 C FOLLOWING STATEMENTS ALLEVIATE THE
51 C PROBLEM.
52 IF(QI .LT. 0.) QI=0.
53 IF(QI .GT. 1./R2*(SSMSR+24./PI)) QI=1./R2*(SSMSR+24./PI)
54 C
55 RETURN
56 END

```



```

1 C
2 C
3 C
4 C
5 C
6 C
7 C
8 C
9 C
10 C
11 C
12 C
13 C
14 C
15 C
16 C
17 C
18 C
19 C
20 C
21 C
22 C
23 C
24 C
25 C
26
27
28 C
29
30 C
31 C
32 C
33 C
34 C
35 C
36 C
37 C
38
39 C
40
41 C
42
43

REAL FUNCTION SOLALT(WT,LAT,DEC)
PURPOSE :
  GIVEN WT - THE HOUR ANGLE IN RADIAN (-180 LE WT LE 180)
  LAT - THE LATITUDE IN RADIAN OF A GIVEN SITE
  DEC - THE SOLAR DECLINATION IN RADIAN FOR A GIVEN DAY
  (-23.45 LE DEC LE 23.45)
THIS ROUTINE WILL CALCULATE THE SOLAR ALTITUDE IN RADIAN
FROM THE TRUE HORIZON.
REQUIRED SUBPROGRAMS : NONE
PROGRAMMING NOTES :
  USAGE - ALT=SOLALT(WT,LAT,DEC)
  PARAMETER TYPES :
    PARAMETERS WT,LAT,DEC ARE SCALARS OF TYPE REAL
    COMMON BLOCKS : NONE

REAL FUNCTION SOLALT(WT,LAT,DEC)
REAL WT,LAT,DEC
SINALT=COS(WT)*COS(LAT)*COS(DEC) + SIN(LAT)*SIN(DEC)
  DUE TO ROUNDING AND/OR CANCELLATION
  ERRORS THE ABOVE STATEMENT MAY ASSIGN
  A VALUE LESS THAN -1 OR GREATER THAN 1
  TO SINALT CAUSING AN ERROR TERMINATION
  IN THE ASIN ROUTINE. THE FOLLOWING
  STATEMENT ELIMINATES THIS PROBLEM.

SINALT=SIGN(AMIN1(ABS(SINALT),1.),SINALT)
SOLALT=ASIN(SINALT)
RETURN
END

```

```

1 C
2 C
3 C
4 C
5 C
6 C
7 C
8 C
9 C
10 C
11 C
12 C
13 C
14 C
15 C
16 C
17 C
18 C
19 C
20 C
21 C
22 C
23 C
24 C
25 C
26
27
28
29
30
31
32
33

REAL FUNCTION RADVEC(D)
PURPOSE :
  GIVEN D - THE SOLAR DECLINATION IN RADIAN
  THIS ROUTINE WILL CALCULATE THE SQUARED RATIO OF
  THE EARTH-SUN DISTANCE TO ITS MEAN FOR THE
  CORRESPONDING DAY.
REQUIRED SUBPROGRAMS : NONE
PROGRAMMING NOTES :
  USAGE - R2=RADVEC(D)
  PARAMETER TYPES :
    PARAMETER D IS A SCALAR OF TYPE REAL
  COMMON BLOCKS -> COMMON/PIE/PI,PI2,PI12,PI180 WHERE
    PI=4*ATAN(1)
    PI2=PI/2
    PI12=PI/12
    PI180=PI/180

REAL FUNCTION RADVEC(D)
COMMON/PIE/PI,PI2,PI12,PI180
REAL D
RADVEC= 0.999847 + 0.001406 * (D/PI180)
RETURN
END

```

```

1000 REAL FUNCTION SOLAZ(WT,LAT,DEC,ALT)
1001
1002 PURPOSE :
1003     GIVEN    WT    - THE HOUR ANGLE IN RADIAN (-180 LE WT LE 180)
1004             LAT    - THE LATITUDE IN RADIAN OF A GIVEN SITE
1005                   (-90 LE LAT LE 90)
1006             DEC    - THE SCLAR DECLINATION IN RADIAN FOR A GIVEN DAY
1007                   (-23.45 LE DEC LE 23.45)
1008             ALT    - THE SOLAR ALTITUDE IN RADIAN FROM THE
1009                   TRUE HORIZON (-90 LE ALT LE 90)
1010
1011 THIS ROUTINE WILL CALCULATE THE SOLAR AZIMUTH IN RADIAN
1012 MEASURED CLOCKWISE FROM SOUTH (0 LE SOLAZ LT 360).
1013
1014 REQUIRED SUBPROGRAMS : NONE
1015
1016 PROGRAMMING NOTES :
1017
1018 USAGE - AZ=SOLAZ(WT,LAT,DEC,ALT)
1019
1020 PARAMETER TYPES :
1021
1022     PARAMETERS WT,LAT,DEC,ALT ARE SCALARS OF TYPE REAL
1023
1024 COMMON BLOCKS - COMMON/PIE/PI,PI2,PI12,PI180 WHERE
1025     PI=4*ATAN(1)
1026     PI2=PI/2
1027     PI12=PI/12
1028     PI180=PI/180
1029
1030 REAL FUNCTION SOLAZ(WT,LAT,DEC,ALT)
1031 REAL WT,LAT,DEC,ALT
1032 COMMON/PIE/PI,PI2,PI12,PI180
1033
1034 ABSWT=ABS(WT)
1035 ABSLAT=ABS(LAT)
1036 ABSDEC=ABS(DEC)
1037 IF (ABS(ALT) .LT. PI2) THEN
1038     SINSAZ=COS(DEC)*SIN(WT)/COS(ALT)
1039
1040     WHEN WT AND ALT SHOULD DIFFER BY + OR - 90
1041     DEGREES BUT DONT DUE TO ROUNDOFF AND/OR
1042     CANCELLATION ERRORS THE ABOVE STATEMENT
1043     MAY ASSIGN A VALUE LESS THAN -1 OR
1044     GREATER THAN +1 TO SINSAZ CAUSING AN
1045     ERROR TERMINATION IN THE ASIN ROUTINE.
1046     THE FOLLOWING STATEMENT ELIMINATES
1047     THIS PROBLEM.
1048
1049     SINSAZ=SIGN(AMIN1(ABS(SINSAZ),1.),SINSAZ)
1050
1051     SAZ=ASIN(SINSAZ)
1052
1053     END IF
1054     IF (ABSLAT .GT. ABSDEC) THEN
1055         HR=ACOS(TAN(DEC)*COTAN(LAT))
1056         IF (ABSWT .LE. HR) THEN

```

```

1057             SOLAZ=SAZ
1058         ELSE
1059             SOLAZ=SIGN(PI,WT) - SAZ
1060         END IF
1061     ELSE
1062         IF (0. .LT. ABSWT .AND. ABSWT .LT. PI) THEN
1063             IF (DEC .EQ. 0.) THEN
1064                 SOLAZ=SIGN(PI2,WT)
1065             ELSE IF (DEC .LT. 0.) THEN
1066                 SOLAZ=SAZ
1067             ELSE
1068                 SOLAZ=SIGN(PI,WT)-SAZ
1069             END IF
1070         ELSE IF (ABSWT .EQ. PI) THEN
1071             IF (ABSLAT .EQ. ABSDEC) THEN
1072                 IF (LAT .GE. 0) THEN
1073                     SOLAZ=PI
1074                 ELSE
1075                     SOLAZ=0.
1076                 END IF
1077             ELSE
1078                 IF (DEC .GT. 0.) THEN
1079                     SOLAZ=PI
1080                 ELSE
1081                     SOLAZ=0.
1082                 END IF
1083             END IF
1084         ELSE
1085             IF (ABSLAT .EQ. ABSDEC) THEN
1086                 IF (LAT .GE. 0.) THEN
1087                     SOLAZ=0.
1088                 ELSE
1089                     SOLAZ=PI
1090                 END IF
1091             ELSE
1092                 IF (DEC .GT. 0.) THEN
1093                     SOLAZ=PI
1094                 ELSE
1095                     SOLAZ=0.
1096                 END IF
1097             END IF
1098         END IF
1099     IF (SOLAZ .LT. 0.) SOLAZ=2.*PI + SOLAZ
1100
1101 RETURN
1102 END

```

C



Rocky
Mountains



Southwest



Great
Plains

U.S. Department of Agriculture
Forest Service

Rocky Mountain Forest and Range Experiment Station

The Rocky Mountain Station is one of eight regional experiment stations, plus the Forest Products Laboratory and the Washington Office Staff, that make up the Forest Service research organization.

RESEARCH FOCUS

Research programs at the Rocky Mountain Station are coordinated with area universities and with other institutions. Many studies are conducted on a cooperative basis to accelerate solutions to problems involving range, water, wildlife and fish habitat, human and community development, timber, recreation, protection, and multiresource evaluation.

RESEARCH LOCATIONS

Research Work Units of the Rocky Mountain Station are operated in cooperation with universities in the following cities:

Albuquerque, New Mexico
Bottineau, North Dakota
Flagstaff, Arizona
Fort Collins, Colorado*
Laramie, Wyoming
Lincoln, Nebraska
Lubbock, Texas
Rapid City, South Dakota
Tempe, Arizona

*Station Headquarters: 240 W. Prospect St., Fort Collins, CO 80526